Data Repository of Paper:
The role of subducted sediments in plate interface dynamics as constrained by Andean forearc (paleo)topography

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DR Section 1. Latitudinal topographic swath profiles of the forearc

Latitudinal swath profiles were performed across straight segments every 0.01° latitude, based on the method described by Hergarten et al. (2014) and with thicknesses such that they cover the entirety of the mapped CC and CD units. In each of these units’ cases, the DEM outside the mapped area was set to Not a Number in order to better isolate their elevations. Swath data values are smoothed by averaging across a 1°-latitude-wide moving window.

DR Section 2. Megathrust seismicity depth parameters

To develop the high spatial resolution data for the six large earthquakes recorded since 1995, we used the finite fault slip distributions of the megathrust earthquakes published by the National Earthquake Information Center (NEIC) of the United States Geological Survey (https://earthquake.usgs.gov/earthquakes/). We projected the inverted slip across the fault on the subduction interface defined by the Slab1.0 model (Hayes et al., 2012) for all earthquakes greater than Mw 7.5 within the Chilean plate boundary. With the projected slip on the subduction interface, we computed the 25th and 75th percentiles of slip-weighted depth distribution over 0.5°-wide latitudinal swaths (see Data Repository Section 3), covering the entire latitudinal extent of our study area. Though a limited record, this computation provides an estimate of where subduction earthquakes tend to concentrate the maximum release of slip along the down-dip extension of the megathrust surface. The identified six events in the catalog correspond to the following earthquakes: (i) the 1995 Mw 8.0 Antofagasta earthquake, (ii) 2007 Mw 7.7 Tocopilla earthquake, (iii) 2010 Mw 8.8 Maule earthquake, (iv) 2014 Mw 8.2 Iquique earthquake and (v) its large Mw 7.7 subduction aftershock, and (vi) the 2015 Mw 8.3 Illapel earthquake (Hayes, 2017) (see Data Repository Figure DR1). Other possible finite slip distributions for those events have been published (e.g., Delouis et al., 2010; Lorito et al., 2011;
Luttrell et al., 2011; Pollitz et al., 2011), including the 1985 Mw 8.0 Valparaiso earthquake (Mendoza et al., 1994), which is not provided by the NEIC catalog. However, we reduced our analysis to the USGS data set, whose fault slip inversions share the same methodology of Ji et al. (2002), so as to maintain consistency.

We then complemented and expanded the analysis of slip depth concentration by looking at the ISC-GEM international seismic catalog (http://www.isc.ac.uk/iscbulletin/search/catalogue/). We selected all the events since 1900 whose centroid locations fall within the slab volume (Hayes et al., 2012) and so are more likely associated to subduction earthquakes. We also discriminate by focal mechanism (when these data are available), selecting only thrust earthquakes.

The surface described by Slab 1.0 was subdivided in several earthquake sub-faults of 0.1º. On each sub-fault, we assigned the closest interface subduction earthquakes and summed all their seismic moments to the particular sub-fault (see Data Repository Figure DR2). For earthquakes larger than the area of the sub-fault, we used the relationship of Allen and Hayes (2017) to determine the source dimensions of length and width. We then distributed uniformly the seismic moment in the sub-faults, such that the center of the uniformly distributed seismic moment of the earthquake was located right on the centroid location given by the ISC-GEM catalogue. We then defined 1º-wide swaths, parallel to the trench, where we computed the seismic-moment-weighted median depth (see Data Repository Section 3). Because of the nature of seismic moment, large earthquakes dominate the statistics. The result is that it is not possible to determine the 25th-75th percentiles as in the previous case of slip distributions for inverted earthquake data.

**DR Section 3. Weighted median and 25th-75th percentiles calculation**

We computed slip and moment-weighted 25th-75th percentiles of depth to evaluate if subduction earthquakes along the studied plate boundary show latitudinal variations of the seismicity depth distribution on the megathrust interface. Using finite slip models from the NEIC catalog and earthquake source parameters from the ISC-GEM catalog we defined 1º-wide latitudinal swaths where we computed the weighted 25th-75th percentiles depth (see DR Section 2). This was accomplished by finding the kth value of depth $d_i$, such that the conditions stated in the following equation are satisfied:

$$\begin{align*}
\sum_{i=1}^{k-1} w_i &\leq \frac{1}{4} ; \frac{1}{2} ; \frac{3}{4} \\
\sum_{i=k}^{n} w_i &\leq \frac{3}{4} ; \frac{1}{2} ; \frac{1}{4}
\end{align*}$$

In this equation, the normalized weight $w_i$ equals the total slip or seismic moment release for each sub-fault within the swath. Because of the nature of seismic moment, large earthquakes dominate the statistics, making it difficult to determine 25th-75th percentiles as with the case of slip distributions for inverted earthquake data.
Supporting figures

**Figure DR1:** Location of finite slip models of large-to-great subduction earthquakes (Mw ≥ 7.5) that occurred in the study area of the Chilean plate boundary since 1995. Slip values are normalized by maximum slip per event. Red curve outlines the trench axis. Models from NEIC catalog (see main text).
Figure DR2: Summed seismic moment of all earthquakes from the ISC-GEM catalog (see main text) projected onto the Slab 1.0 surface (Hayes et al., 2012) over 0.1° sub-faults. At each grid element we assigned the closest interface subduction earthquakes and summed all their seismic moments to the particular sub-fault. Black dashed curve outlines the trench axis.
Figure DR3: Dip (in degrees) of the Nazca-South America plate boundary from the Slab1.0 model (Hayes et al., 2012) down to the -140 km depth contour. Red curve outlines the trench axis.
REFERENCES CITED


