

# Comment on *GSA Today* article by Pavlis et al., 2019: “Subduction Polarity in Ancient Arcs: A Call to Integrate Geology and Geophysics to Decipher the Mesozoic Tectonic History of the Northern Cordillera of North America”

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Pavlis et al. (2019) (abbreviated “P2019”) assert that “geologic and geophysical interpretations lead to fundamentally different conclusions regarding the polarity of subduction along the Cordilleran margin during late Mesozoic time” (p. 1). Their paper is a call to defend a model of uninterrupted eastward subduction beneath continental North America (which we refer to as an “always-Andean” style model) from purportedly contradictory geophysical observations. Our own work, critically cited 12 times on five pages, shows that no such contradiction exists. Neither geology nor geophysics supports always-Andean style subduction since 200 Ma (Sigloch and Mihalynuk 2013, 2017 [abbreviated “SM12013” and “SM2017”]). Instead, both record a Jura-Cretaceous period of simultaneous eastward and westward subduction under a vast archipelago in the northeastern proto-Pacific, analogous to today’s southwestern Pacific. Welded into westward-subducting lithosphere, North America (NA<sub>m</sub>) was pulled into the Archipelago and diachronously overrode it from ca. 155–50 Ma, accreting its arcs and microcontinents.

P2019 portray our Archipelago model as featuring *only* westward subduction, then dismiss it by pointing to the Chugach subduction complex of Alaska, clearly derived from eastward subduction. In reality, our Archipelago model features as much eastward subduction as the Andean-style model, just located further west. Its very essence, missed by P2019, is long-lived subduction of two mature oceans beneath the Archipelago from opposite sides. Our Figure 1A corrects P2019’s rendering of our model (in their fig. 1 and discussion). Our Figure 1B develops today’s Southwest Pacific archipelago as a close tectonic analogue.

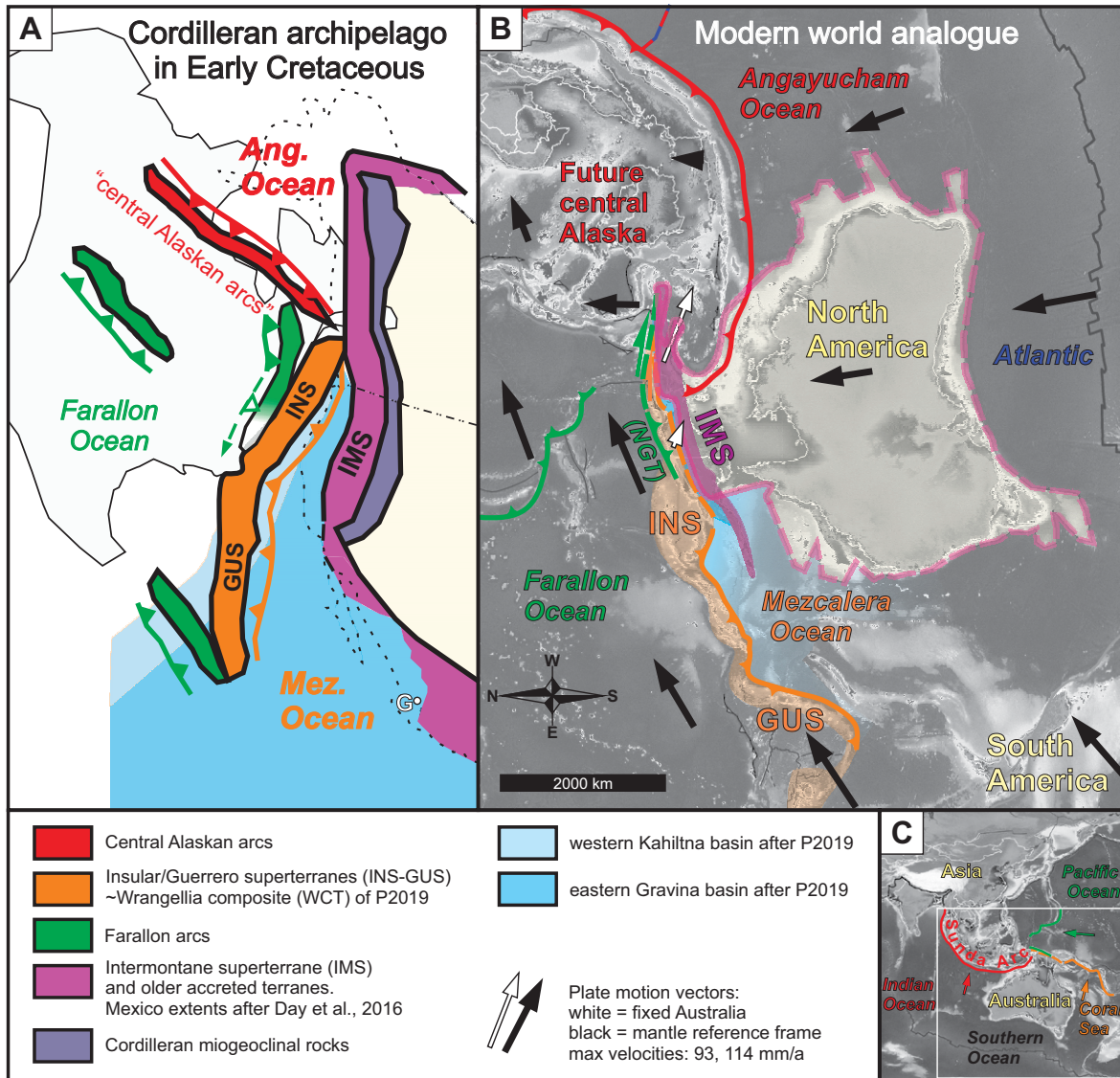
The geologic crux of distinguishing between “always Andean” and “Archipelago” models is the fragmentary record of a suture between NA<sub>m</sub> and the Insular microcontinent (INS; ~“Wrangellia Composite Terrane”/WCT of P2019). Heavy overprinting by the Coast Plutonic Complex results in observational gaps and interpretational ambiguities that are perilously discounted by P2019. Both inboard (continentward) and outboard of this overprinted suture, geologic records support Andean-style subduction over two periods—the “Native” Jurassic–Triassic arc on the cratonic southwestern United States (Dickinson, 2004; Barth et al., 2011; Riggs et al., 2016), and the Cenozoic Cascadia arc in the Pacific Northwest. The “always Andean” model originated as the simplest interpolation between these Jurassic and Cenozoic arcs, with no hypothesis-driven incentive to question Cretaceous times.

Geophysical slab interpretations, including our own (Sigloch et al., 2008; Sigloch, 2011), followed geology’s lead in adopting the “always-Andean margin” interpretation up until SM2013. Geologists’ faith in the validity of this model crossed the community divide as a certainty cast in stone. Dissent among geologists in favor of Archipelago-like models (Moores, 1970, 1998; Schweickert and Cowan, 1975; Ingersoll and Schweickert, 1986; Cowan, 1994; Ingersoll, 2008) did not rise to audible levels for geophysicists. Also, as long as only the supermassive slab walls under NA<sub>m</sub>’s eastern seaboard were clearly visible to seismic tomography, they *had to be* interpreted as Farallon lithosphere in order to match the surviving seafloor isochron record (Atwater, 1989; Engebretson et al., 1985). Even a mantle reference frame was constructed to fix the locus of slab deposition to the continental

margin (van der Meer et al., 2010; used by P2019 in circular reasoning to support their call). An easterly “Farallon slab” was only seriously questioned once a second set of lower-mantle slabs, more westerly and this time truly Farallon, came into focus under western NA<sub>m</sub>, thanks to USArray (Sigloch et al., 2008).

P2019 exaggerate the geophysical uncertainties. Two decades after the iconic Farallon slab images of Grand et al. (1997) on the *GSA Today* cover, not many geophysicists will respond to P2019’s call to challenge “the hypothesis that the deep anomalies are indeed subduction zone remnants” (p. 5). P2019 claim serious problems with slab interpretation because, “These slabs are now in the mantle more than 3000 km from their presumed paleotrench. To restore the pathway over this distance requires multiple assumptions, including the nature of the mantle anomaly, uncertainties in slab sinking rates, and models of absolute plate motion” (p. 2). Confusingly they fail to mention that all these problems (serious indeed) arise only in the “always-Andean” slab interpretation.

Our “tomotectonic” working hypothesis is that a slab indicates the former existence of a trench overhead, because accumulations of lithosphere sink essentially vertically at roughly the same rate everywhere—as expected intuitively and supported by simulations (e.g., Steinberger et al., 2012). Massive slab “walls” observed in the lower mantle under NA<sub>m</sub> must track paleo-trench lines of long-lived, stationary subduction zones (Fig. 1A). Mantle convection under North America, implied to be highly variable over time and space by the “always-Andean” model, becomes simple in the Archipelago framework. Quantitative plate



**Figure 1.** Cordilleran Archipelago in Early Cretaceous times, and its modern-day analogue. (A) Reconstruction of the intra-oceanic Archipelago that was overridden and accreted by North America (NAM) between ca. 150 Ma to ca. 50 Ma. Figure is corrected from Figure 1C in Pavlis et al. (2019), who attempted to render the model of Sigloch and Mihalyuk (2017) but omitted all eastward (Farallon) subduction. Present-day location of North America (NAM) is shaded gray; location in early Cretaceous times (ca. 140 Ma) is shaded yellow. Barbed lines are paleo-trench lines inferred from the presence of massive slab walls in the lower mantle; colored patches are associated arc terranes. Green—Farallon trenches and arcs; orange—Mezcatera trenches and arcs, including Insular Superterrane (INS) and Guerrero Superterrane (GUS); red—Angayucham (Ang) trenches and arcs (future Central Alaska). The Mezcatera (Mez.) [Angayucham] Ocean occupies the space between NAM and the Mezcatera [Angayucham] trench. NAM has started to override the archipelago, flipping orange, intra-oceanic, westward subduction to green, “Andean-style,” eastward subduction. Intermontane Superterrane (IMS) is a pre-accreted part of the NAM margin. (B) The modern-day southwest Pacific represents a closely analogous archipelago (after Sigloch and Mihalyuk, 2017). Base map was obtained by rotating and mirroring the present-day geography; see compass rose and compare to panel (C). Tectonic elements are colored and labeled to match their Cordilleran counterparts in (A). Dashed purple line, filled pale yellow, delineates Australian continental lithosphere. Its position relative to the archipelago is most analogous to NAM’s position ca. 120 Ma, shortly after starting to override the archipelago. Subduction polarity has started to flip from orange to green along Australia’s New Guinea trench segment (NGT). (C) Present-day southwest Pacific. Modern oceans and trenches are colored analogous to panels (A) and (B).

reconstructions permit positioning of NAM relative to these paleo-trenches, and result in a Jura-Cretaceous NAM position well east of the easternmost trenches (Fig. 1A).

NAM in Figure 1A (or Australia in Fig. 1B) first gets pulled by archipelago-ward subduction, but the polarity must flip to continent-ward after arc override. We agree with P2019 that discriminating geologically from

always—continent-ward subduction hinges on establishing the lifetime of the paleo-ocean basin between INS and NAM, and the timing and polarity of its closure. Slab wall locations and distances to Pangean NAM indicate that the basin was much broader than envisaged by P2019 (blue in Fig. 1A); we name it “Mezcatera Ocean,” following Dickinson and Lawton (2001). Slab locations

combined with plate tectonic rules indicate that this Mezcatera Ocean closed by westward subduction under the Archipelago’s eastern border, with absolute trench locations delineated by the eastern-seaboard slab wall now holding this Mezcatera lithosphere. The associated arc must have been built on the older INS micro-continent (Fig. 1A), hence equated with Talkeetna arc—in

contention with P2019, who attribute Talkeetna arc to eastward subduction.

P2019 falsely claim that SM2017 “ignored or dismissed a fundamental observation; namely, that there is compelling geologic evidence that subduction along the northern Cordilleran margin has been east-dipping for at least the last ~125 m.y., and likely can be traced ~75 m.y. further back into the Late Triassic (p. 1).” Our model admits as much eastward subduction as theirs, and we agree that eastward subduction beneath the continental margin built the “Native arc” from Late Triassic (even Permian in southern NAM) through Early Jurassic times. A corresponding “Native Arc slab” is imaged in the lowermost mantle of today’s central Atlantic (Hosseini et al., 2019; SM2017; termed “Atlantis slab” by van der Meer et al., 2010), vertically beneath the reconstructed Triassic margin of NAM (SM2017). This slab is disconnected from the more westerly, shallower Mezcalera slab wall (rather than sloping up toward it—the prediction of “always-Andean” subduction), supporting our hypothesis that INS remained well offshore Pangean NAM.

Our accounting for Chugach complex formation since ca. 145 Ma by Farallon subduction does *not* differ substantially from P2019. Neither does the genesis of the accretionary Franciscan complex (<123 Ma; Dumitru et al., 2010), its pairing with the Sierra Nevada arc, and relation to the Mezcalera suture, which are discussed by SM2017. Accretionary phases of both subduction complexes started in the eastern green patches of Figure 1A, where NAM first overrode the northern Mezcalera trench and its INS arc (orange), forcing a subduction direction flip and nucleating a new, “Andean-style” arc in early Cretaceous times. (Analogous trench flip is ongoing north of New Guinea; Fig. 1B.) Additionally, the Chugach complex probably formed offshore above the Cascadia Root (CR) slab, a deep and voluminous slab under the northwest U.S. that is linked (via the Pacific isochron record; Atwater, 1989; Engebretson et al., 1985) with eastward subduction of the northernmost Farallon plate, from 180+ Ma to today (Sigloch and Mihalynuk, 2013). Any “CR-Chugach” would have accreted to “INS-Chugach” between ca. 90 Ma and the Eocene, the period over which NAM overrode the CR arc (Pacific Rim arc terrane). Thus at least the Chugach’s younger history is related to CR slab deposition. Its fragmentary pre-145 Ma record may also be paired with Pacific Rim arc terrane—countering P2019’s perceived need to

pair Chugach complex with the Talkeetna arc on INS, which our model attributes to Mezcalera westward subduction.

P2019 claim that “there is virtually no evidence for west-dipping subduction anywhere along the inboard margin of the WCT” (p. 2). This representation relies on magmatic overprinting of the suture by the Coast Belt and acceptance of a hypothetical sinistral fault with ~800 km of offset that was invoked to rationalize Jura-Cretaceous subduction zone relicts inboard of INS (Monger et al., 1994). Relicts may extend to NW Washington where they record 35 m.y. of blueschist metamorphism that ceased between 140 and 136 Ma (Cordova et al., 2019). In fact, SM2017 reviewed observations pertaining to arc polarity along the entire Mezcalera Ocean suture (Figs. 1A, 1B). From Alaska to Mexico, this suture is manifested in a dozen collapsed Jura-Cretaceous basins between INS and previously accreted Intermontane Superterrane (IMS). About half of these basins contain ultramafic rocks that may be mantle relicts. Unravelling the story of these relict basins is hampered by the huge volumes of sediment that normally clog them, the intense folding and thrusting that accompanied final collapse, and their thorough overprinting by the thermal-metamorphic welt of the Coast Plutonic Complex, which extends the length of British Columbia (BC).

None of these suture relicts are currently *interpreted* as formed by westward subduction, but a simple, first-order observation shows that they should be: a 35–70-m.y.-long hiatus of volcanic arc strata east of the suture, with simultaneous arc construction west of the suture, on INS. The clearest evidence of this arc hiatus exists across the expanse of Bowser Basin in northwest BC (detailed in SM2017). In southern BC, it has long been recognized that strata on the NAM margin (IMS) lack any record of subduction-related volcanism between ca. 175 and 105 Ma (Thorkelson and Smith, 1989; sparse intrusions within this age bracket have since been dated, but are of uncertain petrogenesis, mostly related to metamorphic terranes, and/or are dated by homogenized, multigrain fractions; e.g., Erdmer et al., 2002). Discovery of rare 163 Ma tuffs on IMS (Mihalynuk et al., 2016) may reduce this volcanic gap, but the petrogenesis of these tuffs is unknown. Regardless, the fundamental observation of an extended arc hiatus east of the suture holds. In mid-coastal BC, Gehrels et al. (2009) note that plutons younger than 160 Ma apparently do not extend eastward from the Coast Belt

into IMS (Stikine terrane). In southern Alaska, too, scant evidence for arc-type magmatism exists north and east of the sutured basins (Kahiltina-Nutzotin) during the hiatus period (e.g., Wilson et al., 2015); whereas INS, south of the suture basins includes magmatic belts (Hart et al., 2004) of Early Cretaceous age: 145–135 Ma, followed by a major flare-up 118–110 Ma, coeval with suturing of the Nutzotin basin (117–114 Ma; Trop et al., 2019). Hence, the subduction of ocean lithosphere that created accretionary complexes and collapsed basins between INS and IMS was likely beneath INS, a conclusion also reached by Dickinson (2004, p. 28).

We stand by our statement that the Chugach subduction complex should not be paired with the Talkeetna arc because their ages do not match (SM2017; Talkeetna arc formed 207–167 Ma) (Amato et al., 2007) or to 153 Ma (Rioux et al., 2007). This basic requirement is not countered by any of P2019’s secondary reasoning. Remarkably, Day, Pavlis, and Amato (2016) themselves advocated a model that allows for pairing of their Chugach complex with an entirely different arc, one that was part of an island archipelago located ~3000 km to the southeast, with *multiple* subduction zones having both east and west polarities (cf. their fig. 6 vs. our Fig. 1A).

P2019 express “virtually no doubt” (p. 3) about (north)east-dipping subduction beneath Talkeetna arc because of its northward younging and increasingly silicic plutons. In reality, such observations cannot constrain subduction polarity without a priori knowledge of the slab’s position relative to the trench. In southwest BC, Canil et al. (2013) show analogous eastward younging of arc magmatism, and shallowing of exhumation levels, for the Talkeetna-correlative, 202–167 Ma Bonanza arc, and they acknowledge the possibility of trench retreat above a westward-subducting slab.

Our analogue identifies Talkeetna arc with the Coral Sea intra-oceanic trench (Fig. 1B). Chugach’s modern analogue is placed continent-ward of the green incipient New Guinea trench. Coupling with the Pacific plate shears these structures westward, analogous to “BajaBC” northward translations of both Chugach complex and Talkeetna arc, through coupling to the Farallon plate (e.g., Day et al., 2016).

NAM paleogeographies need to be consistent with slab geometries, seafloor isochrons, and geologic observations, formulated in a hypothesis-driven work mode that unlocks the full potential of the geological record.

Now that absolute paleo-positioning is available for Cordilleran subduction zones, it is time to give the place a fresh look.

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